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The Basement of the Deccan Traps and Its Madagascar Connection: Constraints from Xenoliths

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ABSTRACT

Paleogeographic reconstructions of India and Madagascar before their late Cretaceous rifting juxtapose the Antongil Block of Madagascar against the Deccan Traps of India, indicating that the Western Dharwar Craton extends below the Deccan lavas. Some recent studies have suggested that the South Maharashtra Shear Zone along the northern Konkan coast of India limits the northern extent of the Western Dharwar Craton, implying that the craton does not extend below the Deccan Traps, raising a question mark on paleogeographic reconstructions of India and Madagascar. The continuity of the Western Dharwar Craton north of the South Maharashtra Shear Zone below the Deccan Traps—or its lack thereof—is critical for validating tectonic models correlating Madagascar with India. In this study, zircons in tonalitic basement xenoliths hosted in Deccan Trap dykes were dated *in situ*, using the U-Pb isotope system. The data furnish U-Pb ages that define three populations at 2527 ± 6 , 2456 ± 6 , and 2379 ± 9 Ma. The 2527 ± 6 Ma ages correspond to the igneous crystallization of the tonalites, whereas the 2456 ± 6 and 2379 ± 9 Ma ages date metamorphic overprints. The results help to establish for the first time that the basement is a part of the Neoproterozoic granitoid suite of the Western Dharwar Craton, which extends northward up to at least Talvade in central and Kihim beach in the western Deccan. By implication, the South Maharashtra Shear Zone cannot be the northern limit of the Western Dharwar Craton. The granitoids are correlated with the Neoproterozoic felsic intrusions (2.57–2.49) of the Masaola suite in the Antongil Block of Madagascar, supporting the existence of a Neoproterozoic Greater Dharwar Craton comprising the Western Dharwar Craton and the Antongil-Masora Block.

Online enhancements: supplementary table.

Introduction

Paleogeographic reconstructions suggest a close link between India and Madagascar from the late Neoproterozoic/early Paleozoic (McWilliams 1981; Stern 1994; Meert et al. 1995; Blasband et al. 2000; De Wit 2003; Sommer et al. 2003; Veevers 2004; Collins and Pisarevsky 2005; Ratheesh Kumar et al. 2014) until their separation in the late Cretaceous (Storey et al. 1995; Torsvik et al. 1998, 2000; fig. 1). The present-day crystalline basement of Madagascar is made up of Archean to Neo-

proterozoic crustal units with equivalents in India and East Africa (Stern 1994; Unrug 1996; Blasband et al. 2000; Kröner et al. 2000; Veevers 2004; Collins and Pisarevsky 2005). The Antongil-Masora Block of Madagascar can be correlated with the Western Dharwar Craton (WDC) of India on the basis of the coherent Paleo/Mesoarchean and Neoproterozoic evolutionary histories of the two crustal units (e.g., Tucker et al. 1999b, 2011a, 2011b, 2014; Schofield et al. 2010; Rekha et al. 2013, 2014). Together, the two terranes constituted the Greater Dharwar Craton (GDC). According to Collins et al. (2003) and others (Collins and Windley 2002; Cox et al. 2004; Fitzsimons and Hulscher 2005; Collins 2006), the Neoproterozoic Antananarivo Block of present-day

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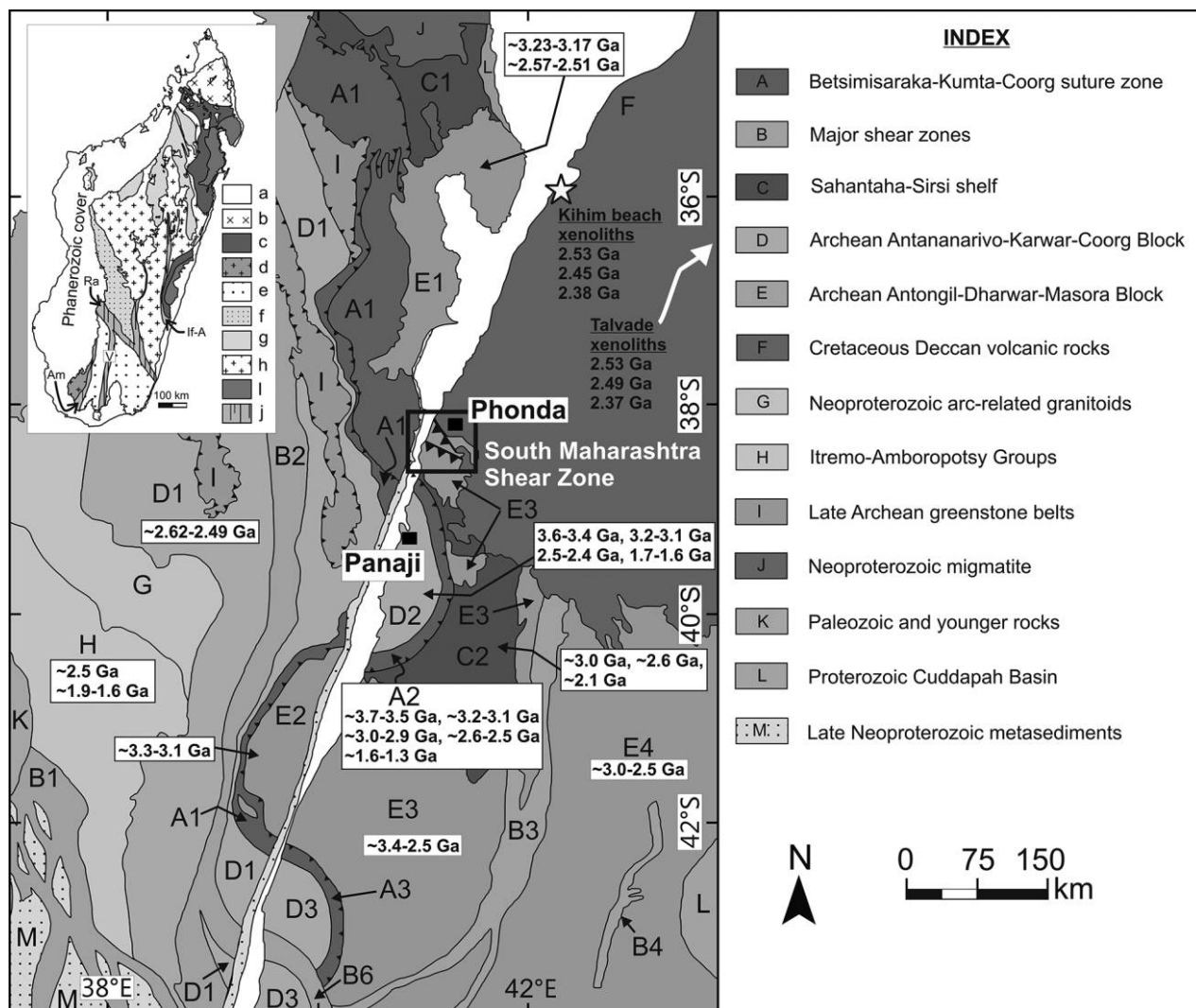


Figure 1. Juxtaposition of India and Madagascar at 120 Ma and suggested correlation of shear zones and tectonic units, as proposed by Ishwar Kumar et al. (2013). The published ages of the crustal units in western India and north-central Madagascar are shown (age data compiled from Balasubramanian et al. 1978; Swami Nath and Ramakrishnan 1981; Naqvi and Rogers 1987; Gupta et al. 1988; Agarwal et al. 1992; Jayananda et al. 1995, 2000; Peaucat et al. 1995; Russel et al. 1996; Tucker et al. 1999b, 2011b, 2014; Chawwick et al. 2000; Paquette et al. 2003; Ghosh et al. 2004; French and Heaman 2010; Schofield et al. 2010; Ishwar Kumar et al. 2013; Rekha et al. 2013). A major part of the Peninsular Indian basement along the Konkan coast is covered and hidden by volcanic rocks of the Deccan Traps, making reliable correlations with Madagascar difficult. The South Maharashtra Shear Zone—which Rekha et al. (2013, 2014) and Rekha and Bhattacharya (2014) claim to represent the northern boundary of the WDC—is marked. The new U-Pb zircon ages from basement xenoliths in Deccan Traps dykes help to establish the northern continuity of the Western Dharwar Craton and its correlation with crustal units in northern Madagascar. Abbreviations: A1, Betsimisaraka suture zone; A2, Kumta suture zone; A3, Coorg suture zone; B1, Ranotsara shear zone; B2, Angavo shear zone; B3, Chitradurga shear zone; B4, Kolar shear zone; B6, Moyar shear zone; C1, Sahantaha shelf; C2, Sirsi shelf; D1, Antananarivo block; D2, Karwar block; D3, Coorg block; E1, Antongil block; E2, Masora block; E3, Western Dharwar Craton; E4, Eastern Dharwar Craton. Inset is an overview geological map of Madagascar showing the major tectonic units. Abbreviations: a, Phanerozoic cover; b, Bemarivo Belt; c, Betsimisaraka Unit; d, Vohibory Unit; e, Androyen Unit; f, Molo Unit; g, Tsaratanana Sheet; h, Antananarivo Block; i, Antongil-Masora Block; j, shear zone systems (Am, Ampanihy; Be, Betsileo; If-A, Ifanadiana-Angavo; Ra, Ranotsara; V, Vorokafotra). A color version of this figure is available online.

Madagascar having East African affinity was amalgamated with India during the early Cambrian collision of the GDC with the Antananarivo Block along the Betsimisaraka Suture. This view is strongly contested by Tucker and coworkers (e.g., Tucker et al. 2011a, 2014), who have questioned the existence of such a suture. These authors propose that the Archean rocks of present-day Madagascar—represented by the Antongil-Masora and Antananarivo domains—were part of the GDC since ca. 3.0 Ga until its fragmentation in the Cretaceous. They also suggest that the Antananarivo Block was amalgamated within the GDC during a Neoproterozoic (2.55–2.48 Ga) accretion event. The Paleoproterozoic Anosy-Androyen unit in southern Madagascar was sutured with the GDC—which comprised the WDC, the Antongil-Masora Block, and the Antananarivo Block—during Ediacaran–early Cambrian time along the Ranotsara-Bongolava shear zone (fig. 1).

A reconstruction of India and Madagascar at ca. 120 Ma (e.g., Katz and Premoli 1979; Collins and Windley 2002; Meert 2003; O'Neill et al. 2003; Ishwar Kumar et al. 2013) just before their rifting juxtaposes the Antongil Block of Madagascar against the late Cretaceous to Paleocene (Lightfoot et al. 1987; Vandamme et al. 1991; Sheth et al. 2001; Mahoney et al. 2002; Pande 2002) Deccan Traps of India (fig. 1). If the Antongil Block and the WDC constituted an expansive GDC until the late Cretaceous, the WDC should extend below the Deccan Trap lavas. Recently, Rekha et al. (2013, 2014) and Rekha and Bhattacharya (2014) have identified a new Paleoproterozoic/Mesoproterozoic shear zone system called the Pernem-Phonda tectonic zone/South Maharashtra Shear Zone between Pernem and Phonda along the northern Konkan coast of western India (fig. 1). These authors claim that the shear zone system is a terrane boundary that limits the northern extent of the WDC. By implication, the WDC does not extend below the Deccan Traps. If true, the Antongil Block would have no counterpart in peninsular India, requiring a relook at the paleogeographic reconstructions of India and Madagascar.

Thus, the continuity of the WDC north of the South Maharashtra Shear Zone—or its lack thereof—is critical for the validation of tectonic models correlating India and Madagascar. Unfortunately, a major part of the Peninsular Indian basement (>500,000 km²; Ray et al. 2008) along the Konkan coast north of the South Maharashtra Shear Zone is covered and hidden by volcanic rocks of the Deccan Traps (figs. 1, 2). As a result, the nature and the geological-geochronological evolution of the basement rocks underlying the Deccan lavas are unconstrained.

In this contribution, we report, for the first time, laser ablation–sector field–inductively coupled plasma mass spectrometer (LA-SF-ICP-MS) U-Pb zircon isotope ages from the rocks derived from the basement below the Deccan Traps. These rocks are exposed as crustal xenoliths in basaltic dykes in the central and western Deccan Traps (fig. 2). The new data help to establish the northern continuity of the WDC and its correlation with crustal units in northern Madagascar.

Regional Geology and Field Setting

The Western Dharwar Craton. The WDC comprises Paleo- to Mesoarchean (3.4–3.0 Ga) tonalite-trondhjemite-granodiorite basement referred to as the Peninsular gneisses (Beckinsale et al. 1980; Taylor et al. 1984; Dhondial et al. 1987; Meen et al. 1992; Peucat et al. 1993; Ishwar Kumar et al. 2013; fig. 2). The basement is overlain by two greenstone-type supracrustal sequences that have been subdivided into the amphibolite to granulite facies Sargur Group of Paleo- to Mesoarchean (3.3–3.1 Ga) age and the greenschist facies Dharwar Supergroup of Neoproterozoic (3.0–2.6 Ga) age (Swami Nath et al. 1976; Janardhan et al. 1979; Nutman et al. 1992; Bouhallier et al. 1993, 1995; Peucat et al. 1995; Ramakrishnan and Vaidyanadhan 2008). The Dharwar Supergroup has been subdivided into two volcano-sedimentary successions, an older Bababudan Group overlain by the younger Chitradurga Group, separated by a disconformity (Swami Nath and Ramakrishnan 1981). The Bababudan Group comprises metamorphosed basalts, layered mafic complexes, mafic/ultramafic schists, and felsic volcanic units together with meta-sedimentary rocks, such as quartzite, phyllite, and banded iron formations. The Chitradurga Group is made up of metasedimentary rocks, such as ortho-quartzite, graywacke, conglomerate, banded iron formation, phyllite, and carbonate interlayered with metavolcanic units. The two volcano-sedimentary successions form the Shimoga and Chitradurga greenstone belts (Swami Nath et al. 1976). The craton underwent a major phase of felsic magmatism accompanied by deformation and metamorphism during the late Neoproterozoic (2.6–2.4 Ga) related to the amalgamation of the Western and Eastern Dharwar Cratons (Friend and Nutman 1991; Chadwick et al. 2000; Moyen et al. 2003; Peucat et al. 2007; Chardon and Jayananda 2008) and the closure of the Goa basin (Rekha et al. 2013). The final phase of igneous activity is represented by the emplacement of mafic dyke swarms at 2.4, 2.0–2.2, and 1.6 Ga (French and Heaman 2010 and references therein).

Recently, Ishwar Kumar et al. (2013) have identified a new arcuate paleo-suture called the Kumta suture zone along the western coast of India. This belt is purported to be the extension of the Betsimisaraka Suture of Madagascar into western India

(fig. 1). The suture separates the Sirsi shelf (Dharwar supracrustals) of the Dharwar cratonic block from the Karwar block (Ishwar Kumar et al. 2013), which comprises tonalite-trondhjemite-granodiorite and amphibolites equivalent to the Peninsular gneisses.

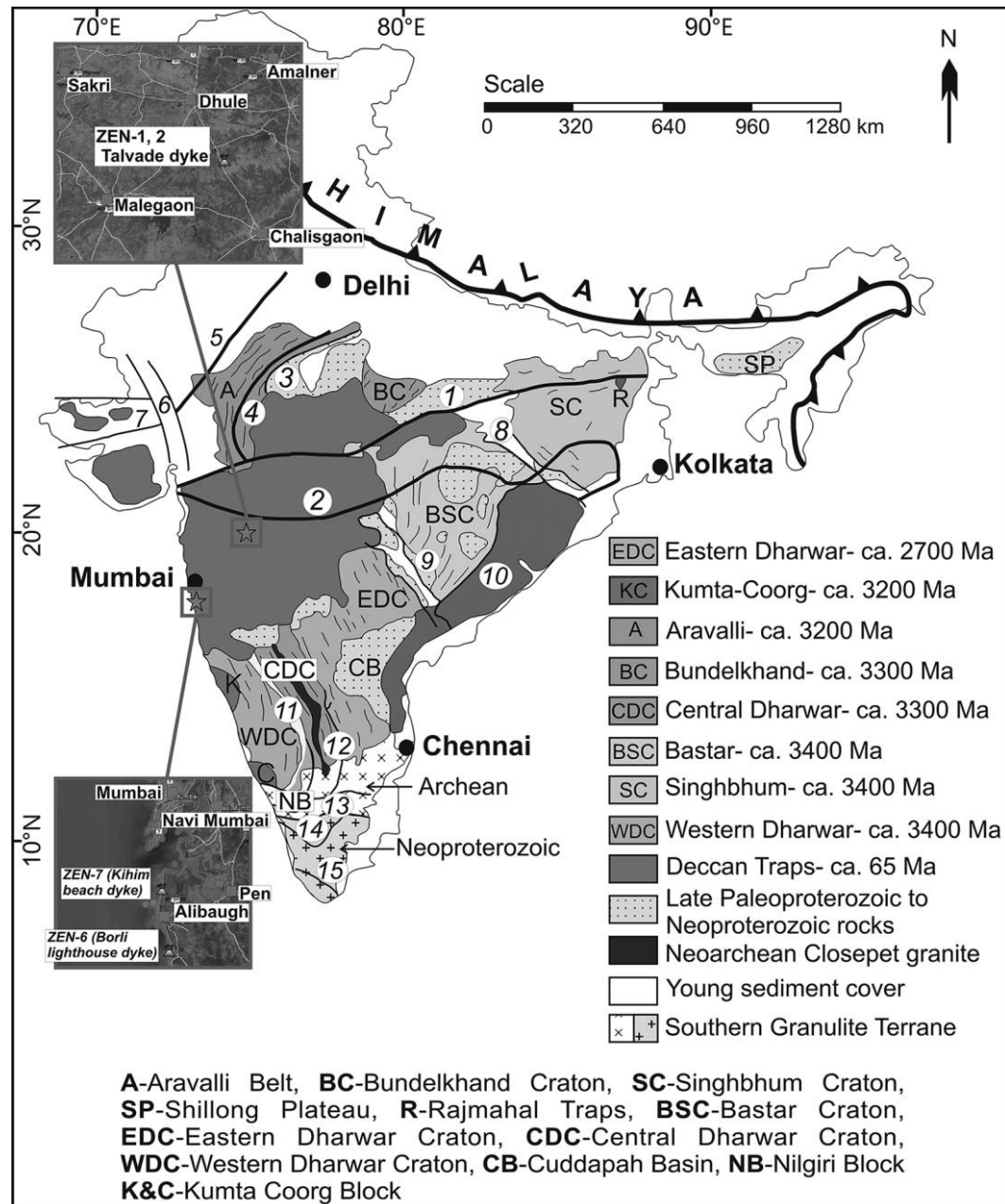


Figure 2. Generalized geological map of peninsular India showing the broad lithotectonic units (modified from Ratheesh Kumar et al. 2014). The Deccan Trap volcanic rocks and the location of Talvade (GPS coordinates: lat 20° 40.24'N, long 74°52.37'E) and Kihim beach (GPS coordinates: lat 18°43.85'N, long 72°51.88'E) xenolith-bearing basaltic dykes are marked with stars. Insets are Google Earth images showing the exact locations of the two dykes. Rifts, shear/suture zones, and mobile belts: 1, Narmada-Son; 2, Satpura Mobile Belt; 3, Great Boundary Fault; 4, Aravalli Fold Belt; 5, Delhi Fold Belt; 6, Cambay Rift; 7, Kutch Rift; 8, Mahanadi Rift; 9, Godavari Rift; 10, Eastern Ghats Mobile Belt; 11, Chitradurga shear zone; 12, Mettur-Kolar shear zone; 13, Palghat-Cauvery shear system; 14, Karur-Kamban-Painavu-Trichur shear zone; 15, Achankovil shear zone. A color version of this figure is available online.

On the basis of the correlation of various shear zones in a reconstruction of India and Madagascar at ~120 Ma, Ishwar Kumar et al. (2013) suggested that the Antananarivo-Karwar-Coorg block (2500–3200 Ma) in the west and the GDC (~2571 Ma) in the east were amalgamated along the Betsimisaraka-Kumta-Coorg suture (fig. 1) diachronously from north to south in the period 1380–750 Ma.

The Deccan Traps and Related Dyke Swarms. The Deccan Traps represent an expansive continental flood basalt province with a composite thickness of ca. 3000 m and covering an area of more than 500,000 km² of the Indian Peninsula (fig. 2). The volcanic rocks are dominated by subaerial flows of tholeiitic basalts with minor occurrences of picrites and alkali basalts (e.g., Beane et al. 1986). The flood basalts were erupted close to the Cretaceous-Tertiary boundary (61–72 Ma; Lightfoot et al. 1987; Vandamme et al. 1991; Sheth et al. 2001; Mahoney et al. 2002; Pande 2002).

The xenolith samples were collected from two E-W-trending dykes at Talvade near Dhule in the central and at Kihim beach near Alibaug in the western part of the Deccan Traps (fig. 2). The Talvade dyke forms a part of the large Narmada-Satpura-Tapi giant dyke swarm of tholeiitic basalts outcropping in the Nandurbar-Dhule region of the central Dec-

can Traps (Sant and Karanth 1990; Ray et al. 2007). The swarm covers an area of ~14,500 km² and contains about 210 dykes of dolerite and basalt intrusive into compound pahoehoe flows of basaltic composition (Ray et al. 2007, 2008). All dykes strike E-W and have been interpreted to be related to the ~N-S regional minimum horizontal compressive stress (σ_3) direction.

The Talvade dyke is 18 km long, ca. 3 m wide, and of basaltic composition (Ray et al. 2008). The dyke contains an assorted variety of xenoliths that are concentrated in the upper part of the dyke along the ridge crest (fig. 3; cf. Ray et al. 2008). The xenoliths are usually angular to tabular and span a wide range of sizes, from millimeter-sized fragments to rafts tens of centimeters across (fig. 3). The rock types include granite gneiss, tonalite gneiss, granite, quartzite, vein quartz, feldspathic granulite, calcareous rock, and fine-grained tuff (this study; Ray et al. 2008). Some of the quartzite and gneissic xenoliths preserve megascopic folds and mylonitic bands (cf. Ray et al. 2008).

The Kihim beach dyke forms a part of the NNW-SSE-trending Konkan dyke swarm outcropping along the Konkan plain (Dessai and Viegas 1995; Ray et al. 2007). The dykes of the Konkan swarm display a large compositional variation from tho-

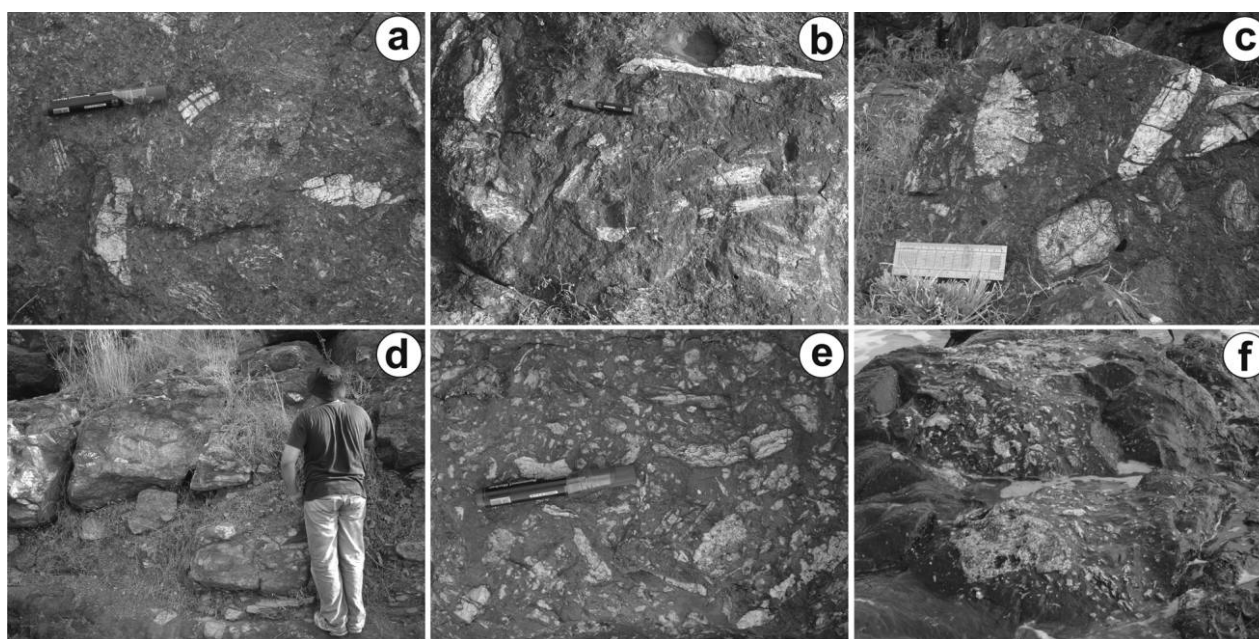


Figure 3. Field photographs documenting the occurrence of an assorted variety of basement xenoliths in basaltic dykes at Talvade (a–e) and Kihim beach (f). The xenoliths are concentrated in the upper central part of the dyke along the ridge crest. They span a wide range of sizes, from millimeter-sized fragments to rafts tens of centimeters across, and include rock types such as granite gneiss, tonalite gneiss, granite, quartzite, vein quartz. Many of the xenoliths are fragmented and heavily altered. A color version of this figure is available online.

leitic to alkalic, ultramafic to silicic and carbonatitic. The Kihim beach dyke is of basaltic composition and contains numerous granitic and tonalitic xenoliths concentrated along the upper central part of the dyke (fig. 3). The size of the xenoliths is variable, from a few millimeters to several centimeters across. In both the Talvade and Kihim beach dykes, basaltic lava has been injected along the foliation/layering of the xenoliths and grain boundaries of the minerals. As a result, many of the xenoliths are fragmented and heavily altered.

North-Central Madagascar. North-Central Madagascar has been divided into five tectonic units (Collins et al. 2000; Kröner et al. 2000; Collins and Windley 2002; Tucker et al. 2011b, 2014; fig. 1): (1) the Antongil-Masora Block, comprising Paleoproterozoic to Mesoproterozoic (~3.2 Ga) gneisses intruded by Neoproterozoic (~2.5 Ga) granites metamorphosed under greenschist facies conditions (Tucker et al. 1999b; Collins et al. 2001; Schofield et al. 2010); (2) the Antananarivo Block, consisting of ~2.5-Ga gneisses interlayered with 820–740-Ma granitoids and gabbros, deformed and metamorphosed at high-grade conditions between 700 and 550 Ma (Tucker et al. 1999b; Kröner et al. 2000; Collins et al. 2003); (3) the Itrimo sheet, comprising Paleoproterozoic to early Neoproterozoic deformed metasedimentary rocks (Cox et al. 1998; Collins et al. 2003) that have been thrust over the Antananarivo block; and (4) the Tsaratanana sheet, consisting of mafic gneisses having an age of 2.8–2.5 Ga (U-Pb zircon) and yielded Meso- to Neoproterozoic xenocrystic zircon and Sm-Nd ages (Tucker et al. 1999b; Collins et al. 2001). The Tsaratanana sheet rocks have been multiply deformed/metamorphosed at ~2.5 Ga and ~0.5 Ga and cut by 800–760 Ma gabbro intrusions (Goncalves et al. 2000); (5) the Bemarivo belt made up of metasediments, granites, and gneisses overlain by deformed metavolcanic rocks metamorphosed under granulite facies conditions at 510–520 Ma (Tucker et al. 1999a). The crustal blocks of central and northern Madagascar are separated from southern Madagascar by the Ranotsara shear zone (Windley et al. 1994).

Analytical Techniques

Zircon grains from granitic/tonalitic xenoliths in the Talvade and Kihim beach dykes were dated in situ in thin sections using a LA-SF-ICP-MS at the Institut für Mineralogie, Westfälische Wilhelms-Universität, Münster. The internal structures of the grains were imaged by cathodoluminescence (CL) and back-scattered electron (BSE) techniques, using a JEOL JSA 6490 secondary electron microscope at the Department of Geology and Geo-

physics, Indian Institute of Technology, Kharagpur. The U-Pb isotope measurements were done on a Thermo-Fisher Scientific Element 2 SF-ICP-MS coupled to a New Wave UP193HE ArF Excimer laser system. The laser was operated at a 10-Hz repetition rate, 5-J/cm² beam energy density, and 12–35-μm spot size. External standardization was done by bracketing groups of 10 unknowns with three measurements of the GJ-1 reference zircon (Jackson et al. 2004). Further details of the analytical and data reduction protocols can be found in Kooijman et al. (2012). Selection of laser ablation spots was based on the internal structure of grain interiors as seen in CL and BSE images. To monitor precision and accuracy, the 91500 reference zircon (Wiedenbeck et al. 1995) was measured ($n = 62$) as unknowns over the course of the analyses and yielded $^{206}\text{Pb}/^{238}\text{U} = 0.179 (\pm 1.5\%, 2\sigma)$ and $^{207}\text{Pb}/^{206}\text{Pb} = 0.0749 (\pm 0.99\%, 2\sigma)$, which match published values within analytical errors. All uncertainties are reported at the 2σ level. The U contents were estimated relative to the GJ-1 reference zircon. Concordia diagrams were constructed using Isoplot 4.15 (Ludwig 2003). The U-Pb isotope ratios and ages are listed in table S1, available online.

Results

Petrographic Description. Four granitoid xenolith samples were selected for the study. Three of them (ZEN-1C, ZEN-2a, ZEN-2c) are from the Talvade dyke, whereas the fourth (ZEN-7a) is from the Kihim beach dyke.

ZEN-1C. This is a quartzofeldspathic rock having two distinct textural domains. One of the domains is strongly foliated, with the foliation plane defined by stretched quartz lenticels showing dynamic recrystallization (fig. 4a). Basaltic melt is seen to have intruded between the foliation planes (fig. 4a) as well as along grain boundaries of minerals (fig. 4d). Evidence of residual strain is seen in the form of undulose extinction in quartz. The dynamically recrystallized quartz grains are overprinted by partial static annealing producing 120° triple junctions. This annealing is possibly related to the heating of the xenolith within the basaltic host magma. The other domain is relatively coarse grained, with coarse-scale layering comprising alternate quartz-rich and plagioclase-rich layers (fig. 4b). The plagioclase is extensively altered and appears cloudy (fig. 4b–4d). Inclusion of iron oxide within grains and along fractures/cracks imparts a dark color to the grains (fig. 4b). The quartz is coarse grained, weakly strained, and equigranular. Euhedral zircons occur within quartz and plagioclase as

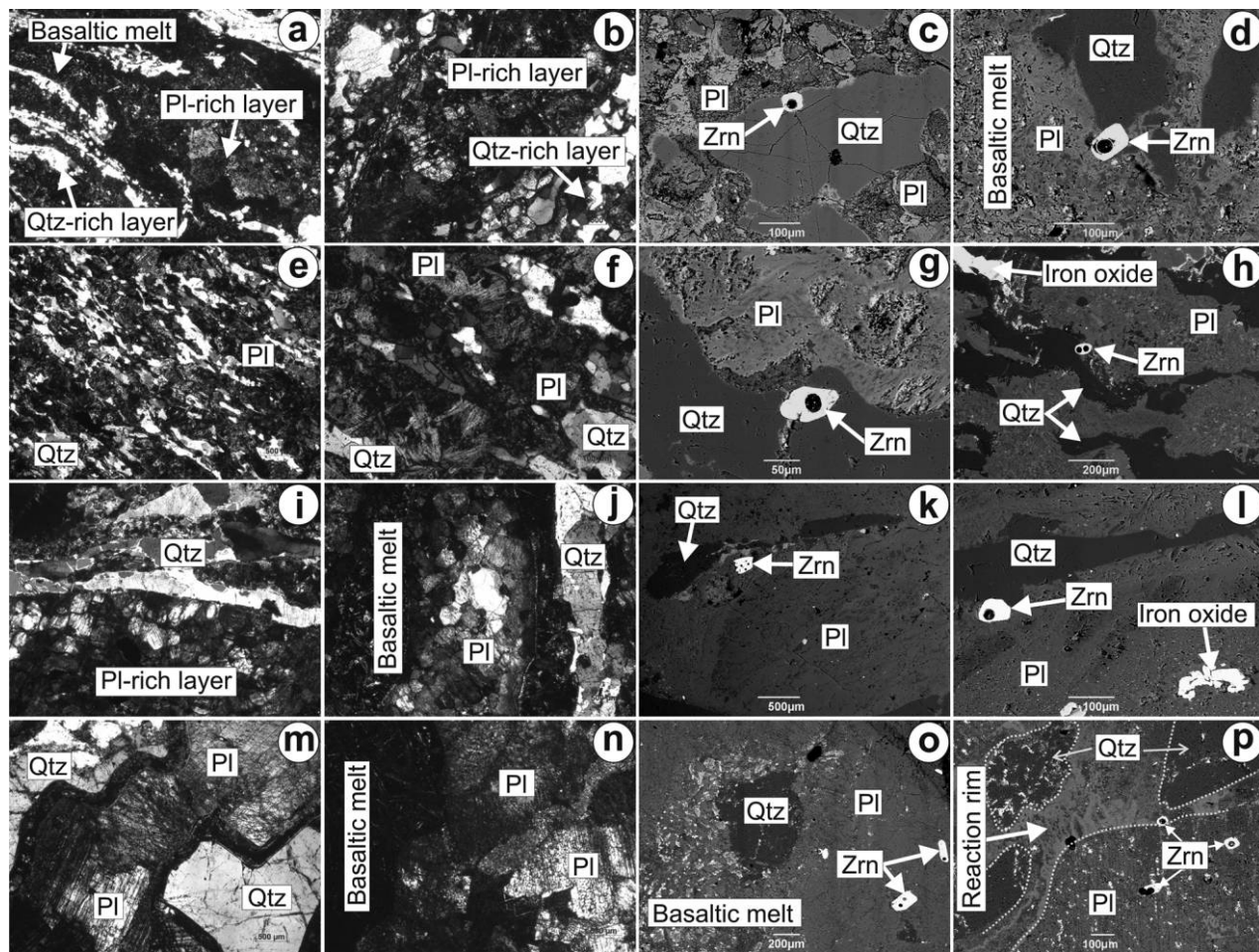


Figure 4. Plane-polarized light photomicrographs (left columns) and BSE images (right columns) illustrating micro-textural relationships in the tonalite/tonalite gneiss basement xenoliths: ZEN-1C (a–d), ZEN-2A (e–h), ZEN-2C (i–l), ZEN-7A (m–p). Extensive alteration of feldspars is observed especially in domains where basaltic melt has intruded the xenoliths. A color version of this figure is available online.

well as along the contact between the two minerals (fig. 4c, 4d). Quartz crystals show evidence for resorption (fig. 4d).

ZEN-2A. The rock is tonalite gneiss with a strong tectonic foliation defined by alternate quartz-rich and feldspar-rich layers (fig. 4e). Quartz grains are stretched to lenticels and show undulose extinction and subgrain formation (fig. 4f). The feldspars are mostly plagioclase and show extensive alteration. They appear clouded and contain opaque (iron oxide) inclusions (fig. 4e–4h). There is extensive iron oxide staining/impregnation along cleavages and grain boundaries (fig. 4f–4h). The quartz has recrystallized to a very fine aggregate and is full of mineral and fluid inclusions. Zircon occurs both in quartz as well as within altered plagioclase.

ZEN-2C. This is relatively coarse-grained tonalite gneiss. The foliation is defined by coarse grains

of elongated quartz alternating with plagioclase-rich layers (fig. 4e). The plagioclase grains show extensive alteration manifested by porosity formation, clouding of the grains, patches of iron oxide inclusions, fluid inclusions, as well as sericite and zoisite inclusions (fig. 4i–4l). The coarse elongated plagioclase laths are recrystallized to a granoblastic mosaic with 120° triple junctions (fig. 4i, 4j). Veins and pockets of basaltic melt intrude the xenolith along foliation. These interfaces are characterized by extensive plagioclase alteration (fig. 4j–4l). Zircon occurs both in quartz as well as within altered plagioclase.

ZEN-7A. This is a very coarse-grained rock comprising mostly plagioclase, with some interstitial quartz. The rock can be classified as a tonalite (fig. 4m). It contains veins and pockets of the host basaltic melt, along which there is extensive alteration of the plagioclase, as in the other samples

(fig. 4n–4p). The altered plagioclase is porous and full of iron oxide inclusions (fig. 4o, 4p). Quartz grains show evidence of resorption (fig. 4o, 4p). A thick reaction rim has developed at the mineral-melt and mineral-mineral interface (fig. 4m, 4p).

Zircon Internal Structure and U-Pb Ages. *Tonalite (ZEN-1C).* Laser ablation analyses of core and rim regions in 16 grains furnish concordant as well as discordant spot ages. The data points define a discordia with an upper intercept age of 2521 ± 10 Ma and a lower intercept age of 90 ± 15 Ma (MSWD = 2; fig. 5). The young lower intercept age indicates alteration/Pb loss of the zircons during the eruption and entrainment of the xenoliths within the Deccan Trap lavas. Therefore, the $^{207}\text{Pb}/^{206}\text{Pb}$ ages from even the discordant data points may represent the real formation ages, especially since most of them cluster within a relatively narrow range. The $^{207}\text{Pb}/^{206}\text{Pb}$ ages define three age populations at 2527 ± 8 , 2487 ± 13 , and 2373 ± 28 Ma. The 2527 ± 8 Ma cluster is identical to the discordia upper intercept age and was measured from domains displaying prominent to relict oscillatory growth zoning (fig. 5b, grains 10, 18, 19, and 20). The 2527 ± 8 Ma

age is therefore interpreted as the emplacement age of the tonalite. Many of the zircons show evidence of fluid/melt-assisted alteration. In BSE images, such grains appear pitted with the development of macroporosity (e.g., grain 3, fig. 5b). In CL images, these zircons show CL-weak alteration patches/layers along cracks/fractures (e.g., grains 3 and 7, fig. 5b). The discordant data points were measured from grains showing these alteration features. Their $^{207}\text{Pb}/^{206}\text{Pb}$ ages define a major age cluster at 2487 ± 13 Ma and a minor one at 2373 ± 28 Ma, both of which are attributed to late Neoproterozoic to early Paleoproterozoic metamorphism of the tonalite.

Tonalite gneiss (ZEN-2A). Most zircon spot analyses give discordant ages that define a discordia indicating near-recent alteration and Pb loss (fig. 6). Thus, the $^{207}\text{Pb}/^{206}\text{Pb}$ ages from even the discordant data points are likely to date real geological events. At the least, they are minimum ages and likely to be geologically significant, since most of them cluster closely together. The $^{207}\text{Pb}/^{206}\text{Pb}$ ages define two populations: one at 2521 ± 21 Ma, measured from U-poor CL-bright domains that preserve prominent-to-faint oscillatory growth zo-

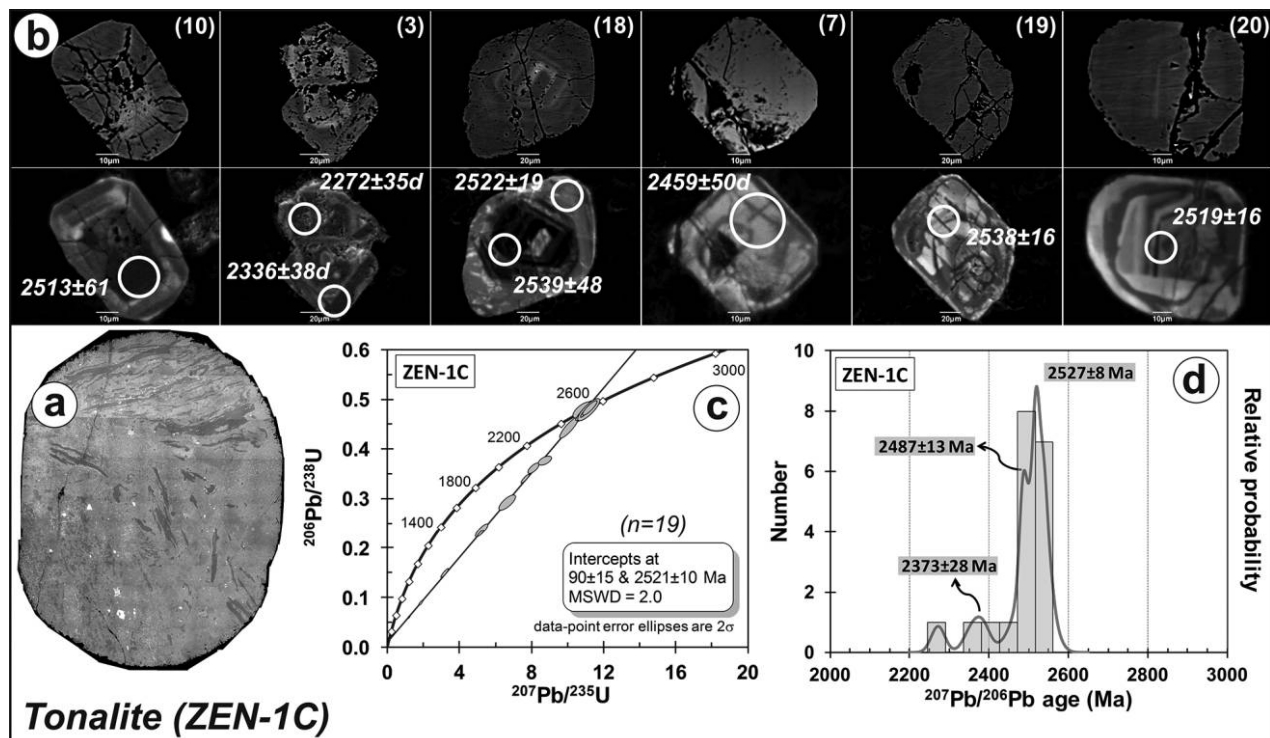


Figure 5. Collage of BSE images showing the overall texture of the tonalite xenolith ZEN-1C (a), BSE and CL images of zircons from the sample with the spot ages marked (d indicates discordant analysis; b), concordia (c), and $^{207}\text{Pb}/^{206}\text{Pb}$ age relative probability density plots (d), respectively, of U-Pb isotope data from zircons. Age populations were computed using the unmix ages option in Isoplot, which uses the unmixing algorithm of Sambridge and Compston (1994). A color version of this figure is available online.

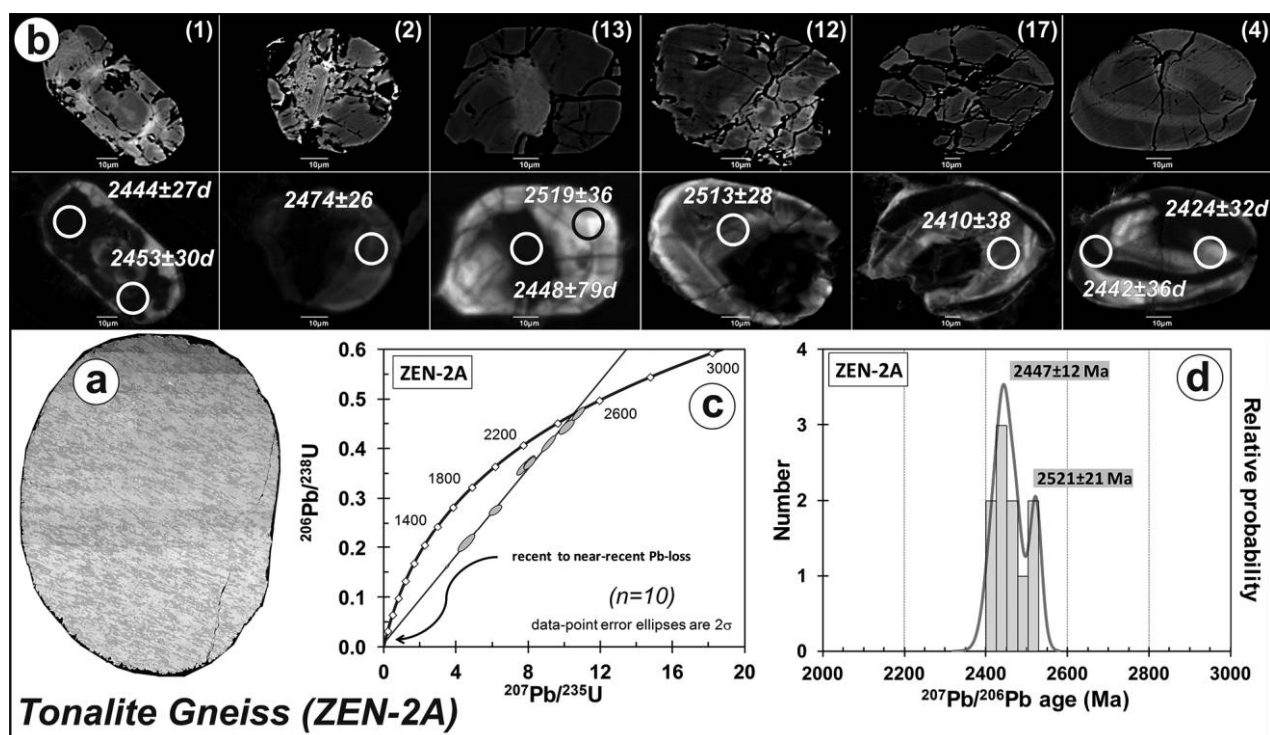


Figure 6. Collage of BSE images showing the overall texture of the tonalite xenolith ZEN-2A (a), BSE and CL images of zircons from the sample with the spot ages marked (d indicates discordant analysis; b), concordia (c), and $^{207}\text{Pb}/^{206}\text{Pb}$ age relative probability density plots (d), respectively, of U-Pb isotope data from zircons. A color version of this figure is available online.

nation (e.g., grains 12, and 13, fig. 6b). This age cluster is identical to the 2527 ± 8 Ma ages obtained from igneous domains in zircons from the tonalite xenolith ZEN-1C and is therefore interpreted to be dating the igneous crystallization of the tonalite. The metamorphic components define a $^{207}\text{Pb}/^{206}\text{Pb}$ age cluster at 2447 ± 12 Ma, similar to the 2487 ± 13 Ma ages obtained from metamorphic zircon in the tonalite ZEN-1C. In CL images, these domains display patchy zoning, have weak CL, and show embayed and irregular resorbed margins (fig. 6b, grains 1, 4, and 13).

Tonalite gneiss (ZEN-2C). Zircon grains in the sample are euhedral and have internal structures typical of igneous zircon. These include oscillatory growth zonation as well as sector zoning (grains 3, 6, 9, 20, and 21, fig. 7b). U-Pb spot ($n = 19$) ages from 12 representative grains define a discordia with upper and lower intercept ages of 2533 ± 9 and 102 ± 25 Ma (MSWD = 0.56), respectively (fig. 7). The young lower discordia intercept indicates recent alteration of the zircons, a feature similar to zircons from the tonalite ZEN-1C and ZEN-2A. The $^{207}\text{Pb}/^{206}\text{Pb}$ ages define a prominent age peak at 2529 ± 9 Ma, analytically indistinguishable from the dis-

cordia upper intercept age. As with the tonalite (ZEN-1C) and tonalite gneiss (ZEN-2A), the 2529 ± 9 Ma age is interpreted to date the emplacement of the magmatic protolith of the gneiss. A minor age peak at 2456 ± 21 Ma is similar to the late Neoproterozoic to early Paleoproterozoic ages from metamorphic zircon in tonalite ZEN-1C and ZEN-2A and is attributed to a metamorphic overprint on the rocks.

Tonalite (ZEN-7A). Most zircon grains show evidence of extensive fluid/melt-induced alteration and plot along a discordia with a young lower intercept age, suggesting recent Pb loss. The $^{207}\text{Pb}/^{206}\text{Pb}$ ages from both concordant and discordant data points define major age peaks at 2445 ± 9 , 2380 ± 10 , and 2313 ± 15 Ma. All the three populations are from altered domains that are pitted and characterized by extensive microporosity in the form of irregular patches or fractures (e.g., BSE images of grains 31, 29, 32, 23, and 35, fig. 8b). The altered regions are bright in BSE images (fig. 8b, grains 31, 19). In some of the grains, core regions are more strongly affected by the alteration, which proceeds inward along irregular fractures (e.g., grains 31, 35, and 19, fig. 8b). Such cores are usually U rich, have

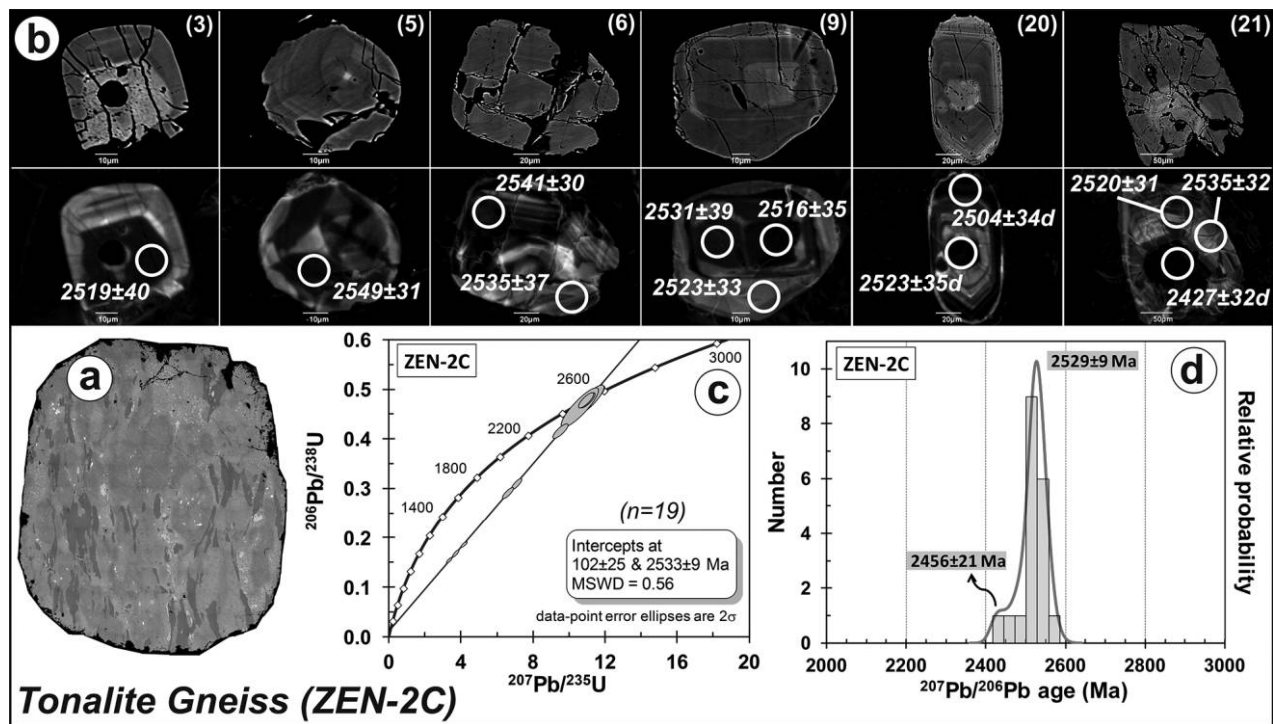


Figure 7. Collage of BSE images showing the overall texture of the tonalite xenolith ZEN-2C (a), BSE and CL images of zircons from the sample with the spot ages marked (d indicates discordant analysis; b), concordia (c), and $^{207}\text{Pb}/^{206}\text{Pb}$ age relative probability density plots (d), respectively, of U-Pb isotope data from zircons. A color version of this figure is available online.

weak CL response, and are possibly metamict. In other grains, the pitted zones define a narrow rim around cores that display patchy CL and BSE zoning (e.g., grains 23 and 29, fig. 8b). The internal structures of the zircon clearly indicate that the three Paleoproterozoic populations (2445 ± 9 , 2380 ± 10 , and 2313 ± 15 Ma) date metamorphic overprints on the rocks. One zircon grain furnished an older age of 2529 ± 39 Ma from a domain having a weak CL growth zoning. In consanguinity with the other tonalite xenoliths, the 2529 ± 39 Ma age dates crystallization of the tonalite. It appears that the metamorphic overprints have almost completely reequilibrated the igneous zircons, accounting for their scarcity.

In summary, the U-Pb isotopic compositions of zircons from the four tonalite/tonalite gneiss xenoliths indicate recent to subrecent Pb loss associated with the emplacement of the Deccan Traps. Igneous components with variably preserved oscillatory growth and sector zoning in all four samples yield a tightly clustered age of 2527 ± 6 Ma (weighted mean, $2\sigma_{\text{internal}}$), which is interpreted as dating late Neoproterozoic tonalitic magmatism in the basement of the Deccan Trap lavas. Zircon domains

that furnish statistically significant age clusters at 2456 ± 6 and 2379 ± 9 Ma (weighted mean, $2\sigma_{\text{internal}}$) are interpreted to date metamorphic overprints on the rocks.

Discussion and Conclusion

Pb Loss in Zircon during Xenolith Entrainment. Zircons in the all the xenolith samples show Pb loss trends, with the discordia lower intercepts scattering between 102 and 56 Ma. The large spread in the lower intercept ages is due to fanning of the discordant analyses from multiple points on the concordia—which date magmatic and metamorphic events—toward the lower intercept. These ages thus might not have any real geological meaning. However, the lower intercepts indicate that the discordance is associated with the disturbance of the U-Pb isotope system in the zircons during the entrainment of the xenoliths in the Deccan Trap basalts.

The Deccan basalts had mantle potential temperature of $\sim 1400^\circ\text{C}$ (Sen 1995, Sen 2001; White and McKenzie 1995) and eruption temperature of $1125^\circ\text{--}1200^\circ\text{C}$ (Melluso et al. 1995). Although the

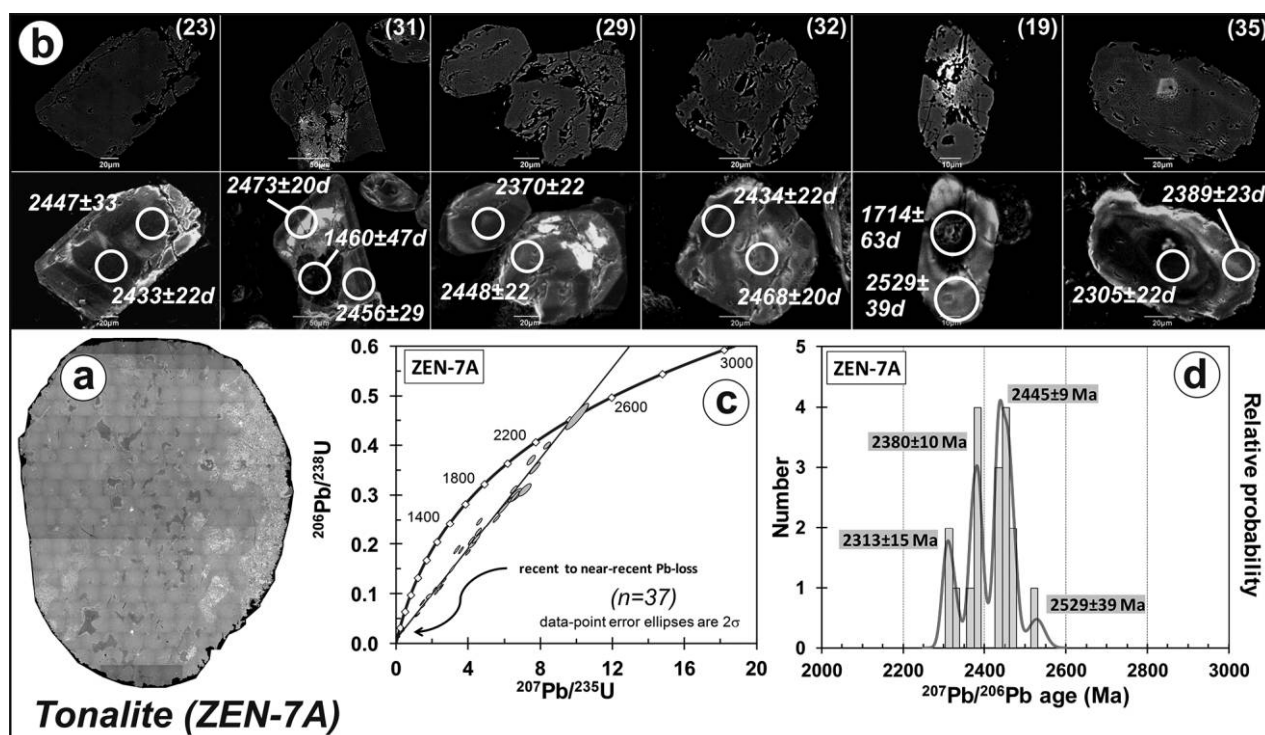


Figure 8. Collage of BSE images showing the overall texture of the tonalite xenolith ZEN-7A (a), BSE and CL images of zircons from the sample with the spot ages marked (*d* indicates discordant analysis; b), concordia (c), and $^{207}\text{Pb}/^{206}\text{Pb}$ age relative probability density plots (d), respectively, of U-Pb isotope data from zircons. A color version of this figure is available online.

general consensus is that the bulk of the eruption took place at 65–66 Ma, existing age data show that magmatism continued long after 61 Ma (Lightfoot et al. 1987; Sheth et al. 2001) to ca. 72 Ma (Mahoney et al. 2002). Thus, the entrained xenoliths may have experienced high-temperature conditions for a long period of time. On the basis of numerical modeling, Bea and Montero (2013) have shown that zircon grains shielded and surviving within xenoliths in mafic magmas at temperatures $>1000^\circ\text{C}$ can experience high temperatures and develop large disturbances in their U-Pb isotopic systematics as a result of enhanced Pb diffusion in a relatively short time, if they are at least partially metamict before entrainment; only zircons with low U concentrations would remain undisturbed by such a thermal overprint (Kooijman et al. 2011). The Pb loss trends exhibited by zircons from the tonalite xenoliths in this study can be explained by a similar process involving high-temperature diffusive Pb loss during the entrainment of the xenoliths in the Deccan basalt magmas. The Pb loss may have been accentuated as a result of the zircons being hosted in feldspars, which have high Pb partitioning and diffusivity (Cherniak 2010; Bea and Montero 2013).

The Northern Extent of the West Dharwar Craton. The Late Neoproterozoic (2.6–2.4 Ga) period in the WDC is marked by a major phase of granitoid magmatism and accompanying deformation-metamorphism. Dhoundial et al. (1987) obtained Neoproterozoic Rb-Sr whole-rock ages from several syn- to post-tectonic granitic plutons in the Goa region (e.g., Chandranath granodiorite/granite: 2650 ± 100 Ma; Dudhsagar granite: 2565 ± 95 Ma; Canacona granite porphyry: 2395 ± 390 Ma). More recently, Ishwar Kumar et al. (2013) have reported U-Pb SHRIMP ages of ~ 2.57 Ga for quartzo-feldspathic gneisses from the Dharwar block. Similar ages (2.6–2.5 Ga, CHIME-monazite) were obtained by Rekha et al. (2013) for the Quepem granitoid in the Goa schist belt. Other parts of the craton also show evidence for extensive Neoproterozoic felsic magmatism. Acid volcanic rocks of the Chitradurga Group (Dharwar Supergroup) give a $^{207}\text{Pb}/^{204}\text{Pb}$ - $^{206}\text{Pb}/^{204}\text{Pb}$ whole-rock isochron age of 2565 ± 28 Ma (Taylor et al. 1984) and an Rb-Sr isochron age of 2520 ± 62 Ma (Bhaskar Rao et al. 1992). The Chitradurga granite pluton has a $^{207}\text{Pb}/^{204}\text{Pb}$ - $^{206}\text{Pb}/^{204}\text{Pb}$ whole-rock isochron age of 2605 ± 18 Ma (Taylor et al. 1984) and a U-Pb zircon age of 2614 ± 10 Ma (Jayananda et al. 2006).

Felsic volcanism and granitic intrusion in the Gadag greenstone belt has been dated at 2588 ± 10 and 2570 ± 10 Ma, respectively (U-Pb zircon; Sarma et al. 2011). The N-S-trending expansive Closepet batholith, which roughly marks the boundary between the western and eastern parts of the Dharwar craton, was also emplaced in the Neoproterozoic (ca. 2.52 Ga; e.g., Friend and Nutman 1991). The large-scale magmatic activity in the craton is broadly synchronous with deformation metamorphism related to the closure of the Goa basin at ca. 2.57 Ga (CHIME monazite; Rekha et al. 2013) and the amalgamation of the Western and Eastern Dharwar Cratons at ca. 2.50 Ga (Friend and Nutman 1991; Chadwick et al. 2000; Moyen et al. 2003; Peucat et al. 2007; Chardon and Jayananda 2008).

The 2.53-Ga tonalitic basement xenoliths within the Deccan Trap dykes at Talvade are thus clearly a part of the expansive Neoproterozoic granitoid suite of the WDC. That the basement beneath the Deccan Traps has a WDC affinity is also supported by the 2.46-Ga ages obtained from metamorphic domains in zircon from the Talvade and Kihim beach tonalite xenoliths. Similar ages, inferred to be dating late Neoproterozoic to early Paleoproterozoic metamorphism, have been reported from metacarbonates at Sandur (~ 2.48 Ga, $^{207}\text{Pb}/^{204}\text{Pb}$ - $^{206}\text{Pb}/^{204}\text{Pb}$ whole-rock isochron; Russell et al. 1996), dolomite-diopside-forsterite marble in the Sargur schist belt (2.48–2.46 Ga, TIMS zircon; Sarangi et al. 2007), and phyllites, schists, and polymict conglomerates from the Goa schist belt (CHIME monazite, 2.44–2.46 Ga; Rekha et al. 2013). The significance of the 2370 ± 22 Ma age from a patchy zircon core in xenolith sample ZEN-7A is not clear. It is tentatively correlated with thermal metamorphism associated with the emplacement of the oldest set of mafic dyke swarm within the Dharwar craton at ca. 2.37 Ga (French et al. 2010).

The results of our study clearly establish the following: (1) the basement of the Deccan Trap lavas comprises tonalitic granitoids that were emplaced at ~ 2.53 Ga and metamorphosed/deformed at ~ 2.46 and ~ 2.37 Ga; (2) the tonalitic basement below the Deccan Traps is a part of the expansive Neoproterozoic granitoid suite of the WDC; (3) the geological evolution of the basement rocks below the Deccan Traps is similar to that of the other parts of the WDC; and (4) the WDC forms the basement to the Deccan Traps lavas, and the cratonic nucleus extends northward up to at least Talvade in the central Deccan and Kihim beach in the western Deccan Traps.

On the basis of CHIME monazite ages, Rekha and Bhattacharya (2014) inferred that the Pernem-

Phonda tectonic zone/South Maharashtra Shear Zone represents a crustal domain formed by long-lived Proterozoic tectonism (~ 2.4 – 1.7 Ga). The authors documented a northward younging of metamorphic episodes across the zone and suggested that the Pernem-Phonda crustal domain was a terrane boundary that limited the northern extent of the WDC. This is not supported by the new zircon age data from the basement xenoliths, which instead show that the WDC extends northward beyond the Pernem-Phonda corridor, thereby necessitating a re-evaluation of the status of the Pernem-Phonda tectonic zone/South Maharashtra Shear Zone.

Paleoproterozoic to Mesoproterozoic ages similar to those for the Pernem-Phonda corridor (e.g., Rekha and Bhattacharya 2014) have also been reported by Ishwar Kumar et al. (2013) from the Kumta suture and the neighboring Karwar and Dharwar blocks. These authors obtained CHIME ages of ~ 1.70 Ga from zircon rims in quartz-phengite-schist from the Kumta suture and K-Ar biotite ages of 1.55–1.64 Ga from the Bondla ultramafic-gabbro complex. They also reported Paleoproterozoic K-Ar biotite ages of ~ 1.73 and 1.75–1.80 Ga from quartzofeldspathic gneisses in the Dharwar block and tonalite-trondhjemite-granodiorites of the Karwar block, respectively. Ishwar Kumar et al. (2013) have interpreted the Paleoproterozoic ages to be dating uplift of the Karwar and Dharwar blocks before their suturing. Thus, the Pernem-Phonda tectonic zone/South Maharashtra Shear Zone could be either (1) intracontinental shear zone within the Dharwar block or (2) suture between the Karwar block and the Dharwar block and therefore equivalent to the Kumta suture.

Correlation with Madagascar. Though the status of the Betsimisaraka suture and its continuation into India as the Kumta suture is still debated, paleogeographic reconstructions juxtaposing Madagascar against India until their late Cretaceous separation is supported by geochronological (e.g., Tucker et al. 1999a, 2011a, 2011b, 2014; Collins and Windley 2002; Collins et al. 2003; Cox et al. 2004; Collins and Pisarevsky 2005; Collins 2006; Schofield et al. 2010; Ishwar Kumar et al. 2013; Rekha et al. 2013, 2014; Rekha and Bhattacharya 2014; Zhou et al. 2015) and geophysical (Chand and Subrahmanyam 2003; Ratheesh Kumar et al. 2014) studies. In a recent study, Ratheesh Kumar et al. (2014) identified zones of conjugate rift-related lithospheric thinning/deformation along the western continental margin of India and the eastern continental margin of Madagascar on the basis of effective elastic thickness of the lithosphere. These zones were inferred to represent the paleo-rift in-

ception points. There is also a close-fit correlation of Moho geometry and bathymetry of the shelf margins for the two continents.

The existence of 2.53-Ga WDC granitoids below the Deccan Traps further adds to the growing body of evidence supporting the existence of a Neoproterozoic GDC comprising the WDC of India and the Antongil-Masora Block of Madagascar. These WDC granitoids can be correlated with the expansive Neoproterozoic felsic intrusives (2.57–2.49 Ga; Tucker et al. 1999b, 2014; Paquette et al. 2003; Schofield et al. 2010) of the Masaola suite in the Antongil Block of Madagascar. Contrary to observations by Rekha et al. (2013) that late Archean intrusive rocks are spatially restricted along the Konkan coast, this study provides evidence for the existence of an expansive suite

of granitoid rocks in the northern part of the WDC similar to those in the Antongil Block. The lack of Neoproterozoic rocks along the Konkan coast is therefore apparent only because the Neoproterozoic basement lies covered by the Deccan Trap volcanic rocks.

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